

Running Title: Effect of CO<sub>2</sub> sequestration on foraminifera

Impact of intentionally injected carbon dioxide hydrate on  
deep-sea benthic foraminiferal survival

Joan M. Bernhard<sup>1</sup>, James P. Barry<sup>2</sup>, Kurt R. Buck<sup>2</sup>, Victoria R. Starczak<sup>3</sup>

<sup>1</sup> Geology and Geophysics Department, MS #52, Woods Hole Oceanographic Institution, Woods Hole, MA 20543 USA

<sup>2</sup> Monterey Bay Aquarium Research Institute, 7700 Sandholdt Road, Moss Landing, CA 95039 USA

<sup>3</sup> Biology Department, MS #50, Woods Hole Oceanographic Institution, Woods Hole, MA 20543 USA

Corresponding Author: Joan M. Bernhard; Tel: 1 508 289 3480; Fax: 1 508 457 2076;

[jbernhard@whoi.edu](mailto:jbernhard@whoi.edu)

**Keywords:** carbon dioxide sequestration, CO<sub>2</sub> injection, climate change, foraminifera, experiment, hypercapnia, meiofauna, Monterey Bay, ocean acidification, protist

## Abstract

Sequestration of carbon dioxide (CO<sub>2</sub>) in the ocean is being considered as a feasible mechanism to mitigate the alarming rate in its atmospheric rise. Little is known, however, about how the resulting hypercapnia and ocean acidification may affect marine fauna. In an effort to understand better the protistan reaction to such an environmental perturbation, the survivorship of benthic foraminifera, which is a prevalent group of protists, was studied in response to deep-sea CO<sub>2</sub> release. The survival response of calcareous, agglutinated, and thecate foraminifera was determined in two experiments at ~3.1 and 3.3 km water depth in Monterey Bay (California, USA). Approximately five weeks after initial seafloor CO<sub>2</sub> release, in situ incubations of the live-dead indicator CellTracker Green were executed within seafloor-emplaced pushcores. Experimental treatments included direct exposure to CO<sub>2</sub> hydrate, two levels of lesser exposure adjacent to CO<sub>2</sub> hydrate, and controls, which were far removed from the CO<sub>2</sub> hydrate release. Results indicate that survivorship rates of agglutinated and thecate foraminifera were not significantly impacted by direct exposure but the survivorship of calcareous foraminifera was significantly lower in direct exposure treatments compared to controls. Observations suggest that, if large scale CO<sub>2</sub> sequestration is enacted on the deep-sea floor, survival of two major groups of this prevalent protistan taxon will likely not be severely impacted, while calcareous foraminifera will face considerable challenges to maintain their benthic populations in areas directly exposed to CO<sub>2</sub> hydrate.

## Introduction

In both the scientific and public arenas, much attention has recently focused on the issues of global warming and climate change. Although debate over the most appropriate mitigation pathway continues, it is certain that the atmospheric concentration of carbon dioxide (CO<sub>2</sub>) has increased significantly in the recent past (e.g., Keeling *et al.*, 1995; Keeling 1998). One option to curtail the rapidly rising atmospheric CO<sub>2</sub> levels under consideration is to sequester waste CO<sub>2</sub> in the deep ocean (e.g., Caldeira, Akai *et al.*, 2005). A number of scenarios have been proposed for such ocean carbon storage, including fertilization of the sea surface with iron to promote phytoplankton growth and accelerate the biological pump, thereby increasing dissolved inorganic carbon (DIC) export to the deep sea (e.g., Buesseler *et al.*, 2004), CO<sub>2</sub> injection at mid-ocean water depths (e.g., Ozaki, 1997), and CO<sub>2</sub> injection onto the deep-sea floor (e.g., Brewer *et al.*, 2000). Ongoing research is attempting to elucidate the benefits and drawbacks of each approach if implemented on the large scale. In particular, impacts on both ocean chemistry and inhabitants must be ascertained, especially given that CO<sub>2</sub> dissolution causes a concomitant decrease in pH (Brewer *et al.*, 2000). Physiological responses to elevated CO<sub>2</sub>, or hypercapnia, and ocean acidification are challenging to organisms in general (reviewed in, e.g., Siebel & Walsh, 2003; Pörtner *et al.*, 2004); it is unclear which taxa, if any, will be unaffected by these environmental pressures.

A series of in situ experiments have been conducted to assess the effects of direct injection of CO<sub>2</sub> on the seafloor, concentrating on the lower bathyal zone since it theoretically provides a longer period of sequestration compared to shallow water (Brewer *et al.*, 1999, 2000; Barry *et al.*, 2005). Although studies have begun to describe the effects of such CO<sub>2</sub> disposal on deep-sea fauna including fish (Tamburri *et al.*, 2000), crustaceans (Barry *et al.*, *submitted*),

echinoderms (Barry *et al.*, *submitted*), and metazoan meiofauna (e.g., Carman *et al.*, 2004; Watanabe *et al.*, 2006; Fleege *et al.*, 2006; Thistle *et al.*, 2007), little is known about the effect of such activities on protista. Because protists comprise a substantial portion of the deep-sea benthos (e.g., Alongi & Pinchon, 1988; Coull *et al.*, 1977; Snider *et al.*, 1984; Gooday *et al.*, 2000; Smith *et al.*, 2002), it is important to establish the effects of bathyal CO<sub>2</sub> release on these single-celled eukaryotes. One study that surveyed the effects of CO<sub>2</sub> release on deep-sea meiobenthos observed that 4.5 weeks after seafloor CO<sub>2</sub> injection many meiofaunal groups (i.e., nematode metazoans and flagellate and amoebae protists) experienced elevated mortality compared to sites removed from CO<sub>2</sub> manipulation (Barry *et al.*, 2004). Robust conclusions about the response of some meiofaunal protists (i.e., allogromiid [thecate] foraminifera and ciliates) to that increased pCO<sub>2</sub> exposure could not be drawn, however, due to low population sizes (Barry *et al.*, 2004) and because an accurate means to determine survival was not implemented for those taxa.

In a continuing effort to better ascertain the response of meiofaunal protists to elevated CO<sub>2</sub> exposure, a study was executed to determine in situ survival response of deep-sea benthic foraminifera to CO<sub>2</sub> release simulating seafloor disposal of this greenhouse gas. For the study, two experiments were conducted, each with a duration of 4.5 weeks in order to evaluate faunal responses to quasi-chronic changes in ocean chemistry, rather than ephemeral changes over short periods (i.e., days). Foraminifera were selected as the study taxon for two main reasons. First, they are a critical link in marine food webs (e.g., Legendre & Le Fèvre, 1995; van Oevelen *et al.*, 2006; Rowe *et al.*, *in press*). Second, while a large proportion of foraminiferal species secrete calcium carbonate shells called tests, the majority lack inorganic tests (so-called allogromiid or thecate forms, e.g., Gooday, 2002) or use detrital particles to construct agglutinated tests (Sen

89 Gupta, 1999), allowing comparative experimentation between calcifying and non-calcifying  
90 species of the same taxonomic group of Rhizarian protists (Adl *et al.*, 2005). This diversity in  
91 foraminiferal test composition is particularly useful when considering the biologic effects of  
92 ocean acidification. It may be hypothesized that the mortality of calcareous foraminiferal species  
93 will be higher in response to deep-sea CO<sub>2</sub> sequestration than the mortality of species of the two  
94 non-carbonate foraminiferan groups (i.e., thecate and agglutinated taxa).

## Materials and Methods

Two replicate experiments were conducted on the seafloor of Monterey Bay, off the California coast (USA). Using the surface support vessel RV *Western Flyer* and the ROV *Tiburon* for each experiment, a set of 40.6-cm diameter PVC cylinders were placed on the seafloor so that ~15 cm extended above the sediment-water interface; 15 cylinders (Experiment 1) and 7 cylinders (Experiment 2) were configured in a circle with a diameter of ~ 20 m. Then, over two to three days, liquid CO<sub>2</sub> was injected into each cylinder using methods described in, e.g., Barry *et al.* (2004), Carman *et al.* (2004), and Fleeger *et al.* (2006). CO<sub>2</sub> hydrate formed immediately as a 'skin' on liquid CO<sub>2</sub> pools. The targeted pH decline within the ~20-m diameter circle compared to in situ pH (~7.8; Thistle *et al.*, 2007 ) was 0.2 pH units because data suggests that pH declines of that magnitude can be an important physiological threshold (e.g., Seibel & Walsh, 2003).

For both experiments, Conductivity-Temperature-Depth instruments (SeaBird Model 19+ CTDs) were deployed on the seabed at locations near the margin, ~3-5 m from the margin, and in the center of the circle of CO<sub>2</sub> cylinders (Fig. 1). Each CTD was equipped with up to 4 Seabird Model SBE 18 pH sensors, positioned from 3 to 50 cm above the seabed. These instruments collected data at ~2 minute intervals throughout each experiment. Perturbations in the pH of surficial sediments were also measured during Experiment 2 within CO<sub>2</sub> cylinders and from 0.1 - 8 m from CO<sub>2</sub> pools, as reported in Barry *et al.* (*submitted*).

During the commencement of Experiment 1, a suite of control samples was obtained, as explained below. The first experiment was initiated from 13-17 December 2004 at a water depth of 3088 m (36.6985°N, 123.0020°W) and terminated from 18-21 January 2005. Experiment 2 was initiated from 12-16 December 2005 at a water depth of 3266 m (35°48.6105'N,

122°34.0952' W) and terminated from 18-21 January 2006. The carbonate chemistry in the region of the experimental sites is typical of the deep Eastern Pacific (WOCE, [www.nodc.noaa.gov/WOCE](http://www.nodc.noaa.gov/WOCE); Sector p17N, St. 10, 38.23733N, 124.93833W, 3087 m depth; DIC=2352.7; TA=2442.3; pH<sub>tot</sub>=7.7834; Omega<sub>CA</sub>=0.93; Omega<sub>AR</sub>=0.6).

As part of each experiment's termination, suites of samples were obtained from a site with no increase in CO<sub>2</sub> (i.e., ~100 m from CO<sub>2</sub> release), from inside the circle of PVC cylinders, and from inside individual PVC cylinders. These three treatments are hereby referred to as Control, Elevated, and Direct Contact, respectively. Elevated treatment samples were further distinguished depending on their distance from the PVC cylinders due to experimental logistics (see below): Elevated Center samples were collected from the approximate middle of the PVC-cylinder circle while Elevated Edge samples were obtained close to PVC cylinders (i.e., within ~1-3 m; Fig. 1). The configuration of core locations was mandated by the fact that this experiment included sampling for multiple additional purposes and by the logistical considerations of working in the deep sea. For example, the ROV manipulator reach and seafloor disturbance in the form of sediment resuspension were critical factors for core placement. More details on the experimental design and concurrent science objectives are presented in Ricketts *et al.* (2005; *submitted*) and Thistle *et al.* (2007).

The fluorogenic probe CellTracker™ Green CMFDA (5-chloromethylfluorescein diacetate; Invitrogen, hereafter referred to as CellTracker Green) was used to distinguish living from dead foraminifera (Bernhard *et al.*, 2006) and, thus, establish survivorship. A fluorogenic probe is a non-fluorescent compound that yields a fluorescent product after structural modification; CellTracker Green accumulates intracellularly in live cells after enzymatic

hydrolysis cleaves the molecule, thereby producing fluorescence (reviewed in Bernhard *et al.*, 1995).

Quantitative sediment samples were obtained using pushcores outfitted with a device to allow injection of concentrated CellTracker Green after pushcore placement into the seafloor. This approach was adopted to allow in situ seafloor incubations with CellTracker Green because temperature and pressure changes during ascent could cause mortality due to sample recovery rather than due to in situ experimental conditions. CellTracker Green is soluble in dimethyl sulfoxide (DMSO); this stock solution was diluted with an equal volume of 0.2  $\mu\text{m}$ -filtered seawater (FSW) to prevent freezing in the injector capillary tubing prior to dispensation.

Initially, the pushcores (7-cm inner diameter) were emplaced into sediments (Fig. 2a) so that the internal header space volume was within a targeted range. To do this, the pushcorer was positioned so that it penetrated the seafloor to a depth placing the sediment-water interface between two core-tube markings. The ROV manipulator arm was then used to carefully squeeze a large diameter (2-cm outer diameter) flexible tube filled with FSW, which flowed through a check valve to displace the concentrated CellTracker Green from the capillary tubing producing a final concentration of  $\sim 1 \mu\text{M}$  CellTracker Green within each core's header space. These cores incubated in situ until the following day, when they were recovered by the ROV. A negative control core, which was injected with DMSO and FSW but not CellTracker Green, was similarly incubated and collected from an area with no increase in  $\text{CO}_2$ , during Experiment 2. To maintain the integrity of the sediment-water interface, the pushcores were not equipped with a stirring device.

After ROV recovery by the support ship, the incubated pushcores were taken within  $\sim 10$  minutes into an environmental room approximating in situ temperature ( $\sim 5^\circ\text{C}$ ). Within an hour,



the overlying water containing CellTracker Green was removed via siphon; the surface 1-cm of each core was sectioned from the underlying sediments, placed in a high density polyethylene (HDPE) container, and fixed in ~3.8% formalin buffered with sodium borate. All Direct Contact treatment cores became disturbed upon ascent due to degassing. In these instances, fine particles were suspended and the sediment-water interface was typically disturbed. Thus, for these cores, the overlying waters were retained in addition to the surface sediment interval or allowed to settle for approximately an hour prior to sectioning.

In the shore-based laboratory, samples were sieved over a 63- $\mu$ m screen with tap water, and the coarser fraction examined with a Leica MZ FLIII stereo dissecting microscope equipped with appropriate epifluorescence optics (excitation 480  $\pm$ 20 nm, emission  $\geq$ 510 nm). Fluorescent foraminifera were isolated, segregated into brightly and dimly fluorescent groups, sorted by species within those groups, and enumerated. Thecate and “soft-shelled” agglutinated taxa, which were not identified to species, were archived in buffered formalin while calcareous and robust agglutinated taxa were air dried and archived on micropaleontology slides. Because a single individual of observed unilocular cylindrical agglutinated foraminiferal taxa (i.e., *Bathysiphon*, *Rhabdammina*, *Rhizammina*) can easily break, those fragments were minimally enumerated to prevent over estimates of population density. Within a sample, it is typically clear which fragments likely arose from the same specimen due to test grain size, diameter, composition, texture, and color. During the picking process, the treatment of each sample (i.e., Control, Elevated, Direct Contact, negative control) was withheld from the microscopist. Species counts were normalized to in situ sample volume to provide abundance estimates.

Abundance data from each experiment were statistically analyzed separately for each taxonomic group. For thecate and agglutinated groups in Experiment 1, a one-way ANOVA was

used to test for differences in density among the Control, Elevated Center, Elevated Edge, and Direct Contact treatments. For Experiment 2, a one-way ANOVA was conducted to test whether mean density of either group differed between the Control, Elevated Edge, and Direct Contact treatments. To homogenize variances, analysis was done on logged transformed thecate data and on square root agglutinated data.

To determine the effects of CO<sub>2</sub> injection on calcareous foraminiferal abundance, the mean abundance in the Control treatment was compared to the mean abundance in the Direct Contract treatment. For this comparison, the mean of the Control treatment was compared to a value of 0 under the null hypothesis with a one sample t-test because calcareous foraminifera were absent in the Direct contact treatment cores. Additionally, a one-way ANOVA was used to test whether mean densities of calcareous foraminifera differed between Control, Elevated Center and Elevated Edge treatments. For Experiment 2, mean calcareous foraminiferal densities in the Elevated Edge and Controls were compared with a two sample t-test and the mean density of the Control treatment was tested in a single sample t-test with a hypothesized mean of zero.

## Results

Only cores that were visibly injected (Fig. 2a) were processed and analyzed. In one case, the capillary tubing parted from a coupling, thus CellTracker Green was not injected into that core so it was excluded from study. Replicate cores from Control and Elevated treatments of each experiment were typically taken within 1 m of each other (Fig. 2b); Direct Contact cores were obtained from different PVC cylinders. Careful manipulation by the ROV pilots resulted in absence of visible sediment-water interface disturbance within the corers, except for those collected from within the PVC cylinders (i.e., Direct Contact cores), which degassed during ascent as noted above.

The negative control core (i.e., injected solely with DMSO, without CellTracker Green) lacked brightly fluorescent foraminifera, but had dimly fluorescent foraminifera. Thus, only brightly fluorescent foraminifera were considered to be living at the time of seafloor incubation with CellTracker Green. In CellTracker Green-incubated cores, the CellTracker Green diffused to at least 1-cm depth because some organisms in the 1-2 cm interval fluoresced brightly.

### *pH changes*

The perturbations of pH measured in bottom waters during both experiments were variable but not large. In an area equivalent to the Elevated Edge treatment, the maximum pH reductions compared to ambient [pH reductions were calculated as perturbations from the background pH measured at the site] values were 0.15 to 0.25 pH units, with average changes less than 0.05 units. In the center of the circle (i.e., Elevated Center treatment), pH reductions sometimes approached 0.1 – 0.2 pH units, but typical pH perturbations were less than 0.05 units.

In areas corresponding to the Elevated Edge treatment, pore-waters of surface sediments at the beginning of Experiment 2 (i.e., 1-3 d after hydrate placement) were ~0.2 units lower than

at control locations (~50 m away) but the pH of Elevated Edge pore-waters at experimental termination did not differ from the pore-water pH of control area sediments (Barry *et al.*, submitted). Changes in the pH of surficial sediment pore waters (i.e., upper ~10 cm) were large within Direct Contact cylinders (~ -2.0 pH units compared to pore-water of control sites; Barry *et al.*, submitted).

#### *Species composition of living foraminiferal assemblages*

The living calcareous foraminiferal assemblage, as determined by bright CellTracker Green labeling, in the Control and Elevated samples was dominated by rotaliid forms including *Uvigerina canariensis*, *Chilostomella oolina*, and *Globobulimina pacifica*; the most common miliolid form was *Quinqueloculina venusta* (Table 1; Fig. 3). The living non-calcareous foraminiferal assemblage in Control and Elevated samples was dominated by thecate and “soft-shelled” agglutinated forms such as saccamminids, as well as *Reophax dentaliniformis*, *R. spiculifer*, *Hormosinella guttifer*, and *Paratrochammina challenger* (Fig. 4). Agglutinated species such as *R. dentaliniformis* and *R. spiculifer* and thecate species dominated Direct Contact samples.

#### *Abundances of live foraminifera*

The foraminiferal abundance between cores within a treatment was highly variable for some treatments but not for others (Table 1). Abundances of all live foraminifera ranged from 4.4 to 5.7 specimens·10cm<sup>3</sup> in Initial Control cores of Experiment 1, and from 6.2 to 11.2 specimens·10cm<sup>3</sup> in Control samples collected at the end of either experiment (Supplemental Data; Fig. 5). Thus, the average abundance of foraminifera in Control samples was 7.7 specimens·10cm<sup>3</sup> (n = 8, SD = 3.0). Foraminiferal abundances within the Elevated samples spanned a wider range (2.6 to 26.0 specimens·10cm<sup>3</sup>;  $\bar{x}$  = 12.1, SD = 8.6) than those within

Direct Contact samples (1.3 to 11.7 specimens·10cm<sup>3</sup>;  $\bar{x}$  = 5.4, SD = 4.0). The average of total abundance of Elevated Edge samples ( $\bar{x}$  = 14.9, SD = 10.4) was higher than all other treatments (Fig. 5a,b).

Abundances of thecate foraminifera averaged 2.3 specimens·10cm<sup>3</sup> (SD = 2.3) in all Control cores and 3.3 specimens·10cm<sup>3</sup> both in Elevated samples and Direct Contact samples (SD = 3.8, 2.5, respectively). Abundances of agglutinated foraminifera averaged 2.1-2.2 specimens·10cm<sup>3</sup> in both the Control and Direct Contact samples, but were nearly twice as abundant in samples collected from within the circle (i.e., Elevated treatments combined; 4.1 specimens·10cm<sup>3</sup>, SD = 3.8). The mean abundances of thecate and agglutinated groups were generally higher in Elevated Edge samples than in other treatments (Fig. 5c-f), the exception being thecate abundances for Experiment 2 (Fig. 5d). The mean density of thecate or agglutinated foraminifera did not differ significantly between the Control, Elevated Center, Elevated Edge or Direct Contact treatments in Experiment 1 (One-way ANOVA, thecate,  $F_{3,7} = 0.39$ ,  $p = 0.762$ ; agglutinates  $F_{3,7} = 0.49$   $p = 0.701$ ) or in Experiment 2 between Control, Elevated Edge, and Direct Contact (One-way ANOVA, thecate,  $F_{2,3} = 0.09$ ,  $p = 0.913$ ; agglutinates,  $F_{2,3} = 0.40$ ,  $p = 0.699$ ).

Calcareous foraminiferal abundances averaged 3.2 specimens·10cm<sup>3</sup> and 4.7 specimens·10cm<sup>3</sup> in Control and Elevated cores, respectively, but no calcareous foraminifera were living in the Direct Contact cores (Table 1; Fig. 5g,h). For Experiment 1, the mean density of calcareous foraminifera did not differ significantly between the Control, Elevated Center or Elevated Edge treatments (One-way ANOVA,  $F_{2,5} = 0.48$ ,  $p = 0.645$ ) and, for Experiment 2, their mean density did not differ significantly between the Control and Elevated Edge treatments (Two sample t-test, -1/x transformed data,  $df = 2$ ,  $t = 2.88$ ,  $p = 0.102$ ). In both experiments, the

270 mean density of calcareous foraminifera in the Control samples was significantly different from  
271 a mean of zero (One sample t-test, Experiment 1:  $t = 6.26$ ,  $df = 2$ ,  $p = 0.025$ , Experiment 2:  
272  $t = 25.96$ ,  $df = 1$ ,  $p = 0.025$ ). Thus, there were significantly more calcareous foraminifera in the  
273 Control treatments than in the Direct Contract treatment, which lacked calcareous foraminifera in  
274 every core of both experiments.

## Discussion

The hypothesis that higher mortality would occur in calcareous foraminifera compared to thecate and agglutinated foraminifera in response to deep-sea CO<sub>2</sub> sequestration is supported by our data. Survivorship of the thecate and agglutinated foraminiferal populations at our Monterey Bay sites were not significantly affected by direct exposure to CO<sub>2</sub> hydrate on the experimental time scale.

Although the water depth at our experimental sites likely exceeded the regional calcite saturation depth and the carbonate compensation depth (CCD), calcareous (calcitic) foraminifera were living in the area as evidenced by Control treatment results. The existence of calcitic foraminifera is not unexpected since most abyssal regions of today's oceans have live calcareous foraminifera (e.g., Bernhard, 1992; Linke & Lutze, 1993). In general, the observed abundances of foraminifera in control samples were comparable to abundances from similar water depths at sites ~150 km to the south (Bernhard, 1992). Using a different viability method (i.e., Adenosine Triphosphate (ATP) assay), abundances of live foraminifera integrated over the surface 1 cm ranged from 0 to ~12 specimens·10cm<sup>3</sup> in water depths from 3319-3728 m ( $\bar{x}$  ~6 specimens·10cm<sup>3</sup>; n = 4; Bernhard, 1992). Deep-sea benthic foraminiferal distributions are known to be patchy on the scale of km to cm (e.g., Bernstein *et al.*, 1978; Bernstein & Meador, 1979), so it is not unusual to document considerable variations in foraminiferal abundance over short distances. Indeed, Bernhard (1992) also noted large variations in foraminiferal abundances between sites, although those cores were separated by kilometers, not meters, as in this study. The species compositions of agglutinated and calcareous assemblages encountered in our samples resemble those previously reported from the region at comparable water depths (e.g., Bernhard, 1992).

Even though some bathyal and abyssal foraminiferal species' abundances exhibit subsurface maxima (e.g., Corliss, 1985; Bernhard, 1992), a down-core analysis of our samples was not feasible. Because, in general, the majority of bathyal foraminifera live in the top cm (e.g., ~52-71%, Gooday, 1986; ~80%, Szarek *et al.*, 2007) and because we expect the maximum physical and chemical changes due to CO<sub>2</sub> hydrate release within the surface sediments, we feel our data reflects the typical survival response of bathyal benthic foraminifera to such environmental perturbations. It is possible that foraminiferans migrated in response to the changing milieu, as observed in other experimental studies (e.g., Alve & Bernhard, 1995; Moodley *et al.*, 1998). In particular, if infaunal specimens migrated upward into the surface cm, the average abundances in treatments affected by CO<sub>2</sub> would have been higher than those of Controls. Indeed, the abundance data for agglutinated taxa were consistently higher, although not significantly different, in Elevated Edge samples compared to Controls (Fig. 5e,f). This trend was not evident for other groups (Fig. 5c,d,g,h). Why foraminifera may have migrated toward potentially stressful concentrations of CO<sub>2</sub> is unclear, except the option of migrating deeper into sediments would have been more stressful for aerobic migrants compared to facultative anaerobes, which are known for foraminifera (e.g., Bernhard & Alve, 1996; Moodley *et al.*, 1998), or if the migrants were incapable of performing complete denitrification (Risgaard-Petersen *et al.*, 2006; Høgslund *et al.*, 2008). Recent findings suggest elevated pCO<sub>2</sub> exposure is stressful to at least some meiofauna because higher numbers of harpacticoid copepods emerged from sediments in response to elevated pCO<sub>2</sub> compared to copepod emergence rates at control sites (Thistle *et al.*, 2007).

Degassing of cores collected from PVC cylinders during ascent may have minimized the observed foraminiferal abundances in the Direct Contact cores. We discount this possibility



because the average density of thecate and agglutinated morphotypes in Direct Contact samples was near or equal to their density in Control cores. If significant abundance dilution occurred due to degassing disturbance, we would expect all three groups of foraminifera to be similarly affected, which was not the case.

Observed patterns and magnitudes of pH change resemble those reported for similar experiments (Barry *et al.*, 2005). Importantly, the observed bottom-water pH decreases in areas corresponding to Elevated Center treatments (~ 0.05 units for Experiment 1) were small compared to those declines corresponding to Elevated Edge treatments (initially 0.2 pH units, but ~ 0.1 to 0.15 pH units at experiment end). Thus, it is not surprising that all three groups of benthic foraminifera (i.e., calcareous, agglutinated, thecate) tolerated the Elevated Center treatments because those sediments apparently did not experience extended large decreases in pH and, by inference, extended large increases in pCO<sub>2</sub>. It is also noteworthy, however, that thecate and agglutinated foraminiferal survival was not negatively impacted by the substantial pH decrease (> 1 pH unit) in the Direct Contact treatments.

Economically viable protocols of deep-sea CO<sub>2</sub> sequestration include formation of deep-sea CO<sub>2</sub> lakes and near-bottom injection of CO<sub>2</sub> (reviewed in Herzog *et al.*, 2000; Caldeira & Akai *et al.*, 2005). In all cases, a plume of CO<sub>2</sub>-rich seawater will lead from liquid CO<sub>2</sub> or CO<sub>2</sub> hydrate drifting with currents and diffuse, eventually being diluted (Caldeira and Akai *et al.*, 2005). Thus, CO<sub>2</sub> and pH gradients will range from substantial (e.g., the lowest pH measured is 4.5, a value measured within cm of CO<sub>2</sub> hydrate; Brewer *et al.*, 2000) to mild (i.e., <-0.05 pH units) at some distance, depending on specific mixing conditions. Our Direct Contact treatment recreates the CO<sub>2</sub> lake approach, although on a much smaller scale. Benthos beneath any CO<sub>2</sub>-hydrate pool will be directly exposed for long periods of time unless they can migrate from the

hydrate pool (Tamburri *et al.*, 2000). Deep-sea foraminifera move very slowly ( $<25 \mu\text{m}\cdot\text{min}^{-1}$ ; Gross, 2000), so active avoidance at these spatial scales is an implausible escape mechanism for them. The pH changes in our experiments are similar to those under any carbon sequestration option: conditions of our Elevated treatments are within the magnitude of pH change expected along the dilution gradient resultant from presently proposed mitigation scenarios.

The in situ response of foraminifera (and any other group of deep-sea organisms) to such  $\text{CO}_2$ -hydrate exposure is unknown over time periods longer than our experiments. Our results indicate that although the calcareous foraminiferal population tolerates short exposures to a 0.2 pH unit decrease, their survival response to direct  $\text{CO}_2$ -hydrate exposure results in a significant increase in mortality. If, however, plumes of dissolved  $\text{CO}_2$  hydrate are diffused from the seafloor over a shorter time scale than our experiments, it is possible that calcareous foraminiferal survival will not be significantly impacted by large-scale  $\text{CO}_2$  sequestration.

Although the effects of localized deep-sea  $\text{CO}_2$  sequestration appear to negatively impact calcareous foraminifera (Ricketts *et al.*, 2005; *submitted*; this study), two other protistan groups (i.e., flagellates and amoebae; Barry *et al.*, 2004), and a number of metazoan meiofauna (Barry *et al.*, 2004; Carman *et al.*, 2004; Watanabe *et al.*, 2006; Fleeger *et al.*, 2006; Thistle *et al.*, 2007), our findings suggest that survival of at least some protistan meiofauna (i.e., thecate and agglutinated foraminifera) are not similarly influenced by direct exposure to  $\text{CO}_2$  hydrate. Because the abundance and diversity of thecate and agglutinated foraminifers are considerable in bathyal and abyssal sediments (e.g., Gooday *et al.*, 1998; 2000; Smith *et al.*, 2002) and their abundance can exceed that of other meiofauna (e.g., Gooday *et al.* 2000), at least one major group of deep-sea meiofauna will likely not collapse if large-scale sequestration of  $\text{CO}_2$  is implemented on the deep-ocean floor. Furthermore, our data showing that thecate and agglutinated

367 foraminiferal abundances do not significantly decline in response to direct exposure to CO<sub>2</sub>  
368 hydrate at these spatial and temporal scales substantiate recent laboratory findings documenting a  
369 shallow-water thecate foraminiferan species survives ~2-week exposure to extremely high pCO<sub>2</sub>  
370 (200,000 ppm) where some specimens even reproduced (Bernhard et al., *in press*).

## Acknowledgements

Appreciation is extended to Eric Pane for synopsis of pore-water pH data; Craig Okuda for designing the push-core injectors; Patrick Whaling, Chris Lovera, Linda Kuhn, Kevin Carman, John Fleeger, David Thistle, and Erin Ricketts for sampling assistance; the Captain and crew of the RV *Western Flyer* and the ROV *Tiburon*; Marti Jeglinski for laboratory assistance; David Smith, Subject Editor and the two anonymous reviewers for their comments on a previous manuscript version. This work was funded by the Monterey Bay Aquarium Research Institute (project 200002; to JPB), US Department of Energy grant # DE-FG02-03ER63696 (to J. P. Kennett and J.M.B.), and NSF OCE-0725966 (to J.M.B.).

## References

- Adl SM, Simpson AGB, Farmer MA, Andersen RA, Anderson OR, Barta JR, Bowser SS, Brugerolle G, Fensome RA, Fredericq S, James TY, Karpov S, Kugrens P, Krug J, Lane CE, Lewis LA, Lodge J, Lynn DH, Mann DG, McCourt RM, Mendoza L, Moestrup O, Mozley-Standridge SE, Nerad, TA, Shearer CA, Smirnov AV, Spiegel FW, Taylor MFJR (2005) The new higher level classification of eukaryotes with emphasis on the taxonomy of protists. *Journal of Eukaryotic Microbiology*, **52**, 399-459.
- Alongi DM, Pinchon M (1988) Bathyal meiobenthos of the western Coral Sea: distribution and abundance in relation to microbial standing stocks and environmental factors. *Deep-Sea Research*, **35**, 491-503.
- Alve E, Bernhard JM (1995) Vertical migratory response of benthic foraminifera to controlled decreasing oxygen concentrations in an experimental mesocosm. *Marine Ecology Progress Series*, **116**, 137-151.
- Barry JP, Buck KR, Lovera CF, Kuhnz L, Whaling PJ, Peltzer ET, Walz P, Brewer PG (2004) Effects of direct ocean CO<sub>2</sub> injection on deep-sea meiofauna. *Journal of Oceanography*, **60**, 759-766.
- Barry JP, Buck KR, Lovera C, Kuhnz L, Whaling PJ (2005) Utility of deep sea CO<sub>2</sub> release experiments in understanding the biology of a high-CO<sub>2</sub> ocean: Effects of hypercapnia on deep sea meiofauna. *Journal of Geophysical Research—Oceans*, **110**, Issue: C9 Article Number: C09S12.
- Barry JP, Buck KR, Lovera C, Brewer PG, Seibel BA, Drazen JC, Tamburri MN, Whaling PJ, Kuhnz L, Pane E (*submitted*) Sensitivity of deep-sea animals to a high-CO<sub>2</sub> ocean. *Marine Ecology Progress Series*.

403 Bernhard JM (1992) Benthic foraminiferal distribution and biomass related to pore-water  
 404 oxygen content: Central California Continental Slope and Rise. *Deep-Sea Research*, **39**:  
 405 585-605.

406 Bernhard JM, Alve E (1996) Survival, ATP pool, and ultrastructural characterization of benthic  
 407 foraminifera from Drammensfjord (Norway): response to anoxia. *Marine*  
 408 *Micropaleontology*, **28**, 5-17.

409 Bernhard JM, Newkirk SG, Bowser SS (1995) Towards a non-terminal viability assay for  
 410 foraminiferan protists. *Journal of Eukaryotic Microbiology*, **42**, 357-367.

411 Bernhard JM, Ostermann DR, Williams DS, Blanks JK (2006) Comparison of two methods to  
 412 identify live benthic foraminifera: a test between Rose Bengal and CellTracker Green  
 413 with implications for stable isotope paleoreconstructions. *Paleoceanography*, **21**,  
 414 PA4210, doi:10.1029/2006PA001290.

415 Bernhard JM, Mollo-Christensen E, Eisenkolb N, Starczak VR (*in press*) Tolerance of  
 416 allogromiid foraminifera to severely elevated carbon dioxide concentrations: Implications  
 417 to future ecosystem functioning and paleoceanographic interpretations. *Global and*  
 418 *Planetary Change*.

419 Bernstein BB, Meador JP (1979) Temporal persistence of biological patch structure in an  
 420 abyssal benthic community. *Marine Biology*, **51**, 179-183.

421 Bernstein BB, Hessler RR, Smith CR, Jumars PA (1978) Spatial distribution of benthic  
 422 foraminifera in the abyssal central North Pacific. *Limnology and Oceanography*, **23**,  
 423 401-416.

424 Brewer PG, Friederich G, Peltzer ET, Orr FM (1999) Direct experiments on the ocean disposal  
 425 of fossil fuel CO<sub>2</sub>. *Science*, **284**, 943-945.

426 Brewer PG, Peltzer ET, Friederich G, Aya I, Yamane K (2000) Experiments on the ocean  
 427 sequestration of fossil fuel CO<sub>2</sub>: pH measurements and hydrate formation. *Marine*  
 428 *Chemistry*, **72**, 83-93.

429 Buesseler KO, Andrews JE, Pike SM, Charette MA (2004) The effects of iron fertilization on  
 430 carbon sequestration in the Southern Ocean. *Science*, 304, 414-417.

431 Caldeira, K, Akai M, *et al.* (2005) Ocean Storage. In: *IPCC Special Report on Carbon Dioxide*  
 432 *Capture and Storage*, B. Metz, O. Davidson, H. de Coninck, M. Loos, L. Meyer (Eds),  
 433 Cambridge U. Press, Cambridge.

434 Carman KR, Thistle D, Fleeger JW, Barry JP (2004) Influence of introduced CO<sub>2</sub> on deep-sea  
 435 metazoan meiofauna. *Journal of Oceanography*, **60**, 767-772.

436 Corliss BH (1985) Microhabitats of benthic foraminifera within deep-sea sediments. *Nature*, **314**,  
 437 435-438.

438 Coull BC, Ellison RL, Fleeger JW, Higgins RP, Hope WD, Hummon WD, Rieger RM, Sterrer  
 439 WE, Thiel H, Tietjen JH (1977) Quantitative estimates of the meiofauna from the deep  
 440 sea off North Carolina, USA. *Marine Biology*, **39**, 233-240.

441 Fleeger JW, Carman KR, Welsenhorn PB, Sofranko H, Marshall T, Thistle D, Barry JP (2006)  
 442 Simulated sequestration of anthropogenic carbon dioxide at a deep-sea site: Effects on  
 443 nematode abundance and biovolume. *Deep-Sea Research I*, **53**, 1135-1147.

444 Gooday AJ (1986) Meiofaunal foraminiferans from the bathyal Porcupine Seabight (northeast  
 445 Atlantic): size structure, standing stock, taxonomic composition, species diversity and  
 446 vertical distribution in the sediment. *Deep-Sea Research*, **33**, 1345-1373.

447 Gooday AJ (2002) Organic-walled allogromiids: aspects of their occurrence, diversity and  
 448 ecology in marine habitats. *Journal of Foraminiferal Research*, **32**, 384-399.

449 Gooday AJ, Bett BJ, Shires R, Lambshead PJD (1998) Deep-sea benthic foraminiferal species  
 450 diversity in the NE Atlantic and NW Arabian sea: a synthesis. *Deep-Sea Research II*, **45**,  
 451 165-201.

452 Gooday AJ, Bernhard JM, Levin LA, Suhr S (2000) Foraminifera in the Arabian Sea oxygen  
 453 minimum zone and other oxygen-deficient settings: taxonomic composition, diversity,  
 454 and relation to metazoan faunas. *Deep-Sea Research II*, **47**, 25-54.

455 Gross O (2000) Influence of temperature, oxygen and food availability on the migrational  
 456 activity of bathyal benthic foraminifera: evidence by microcosm experiments.  
 457 *Hydrobiologia*, **426**, 123-137.

458 Herzog H, Eliasson B, Kaarstad O (2000) Capturing greenhouse gases. *Scientific American*, **282**,  
 459 72-79.

460 Høglund S, Revsbech NP, Cedhagen T, Nielsen LP, Gallardo VA (2008) Denitrification, nitrate  
 461 turnover, and aerobic respiration by benthic foraminiferans in the oxygen minimum zone  
 462 off Chile. *Journal of Experimental Marine Biology and Ecology*, **359**, 85-91.

463 Keeling CD (1998) Rewards and penalties of monitoring the earth. *Annual Review of Energy*  
 464 *and the Environment*, **23**, 25-82.

465 Keeling CD, Whorf TP, Wahlen M, Van der Plicht J (1995) Interannual extremes in the rate of  
 466 rise of atmospheric carbon dioxide since 1980. *Nature*, **375**, 666-670.

467 Legendre L, Le Fèvre J (1995) Microbial food webs and the export of biogenic carbon in oceans.  
 468 *Aquatic Microbial Ecology*, **9**, 69-77.

469 Linke P, Lutze GF (1993) Microhabitat preferences of benthic foraminifera—a static concept or  
 470 a dynamic adaptation to optimize food acquisition? *Marine Micropaleontology*, **20**, 215-  
 471 234.



472 Moodley L, van der Zwaan GJ, Rutten GMW, Boom RCE, Kempers AJ (1998) Subsurface  
 473 activity of benthic foraminifera in relation to porewater oxygen content: laboratory  
 474 experiments. *Marine Micropaleontology*, **34**, 91-106.

475 Ozaki M (1997) CO<sub>2</sub> injection and dispersion in mid-ocean depth by moving ship. *Waste*  
 476 *Management*, **17**, 369-373.

477 Pörtner HO, Langenbuch M, Reipschläger A (2004) Biological impact of elevated ocean CO<sub>2</sub>  
 478 concentrations: Lessons from animal physiology and earth history. *Journal of*  
 479 *Oceanography*, **60**, 705-718.

480 Ricketts ER, Kennett JP, Hill TM, Barry JP (2005) Effects of CO<sub>2</sub> hydrate on deep-sea  
 481 foraminiferal assemblages. Proceedings of the Fifth International Conference on Gas  
 482 Hydrates, Trondheim, Norway, 3, (3020): 839-847.

483 Ricketts ER, Kennett JP, Hill TM, Barry JP (*submitted*) Effects of CO<sub>2</sub> hydrate emplacement on  
 484 deep-sea foraminiferal assemblages: 3600 m on the California Margin. *Marine*  
 485 *Micropaleontology*.

486 Risgaard-Petersen N, Langezaal AM, Ingvardsen S, *et al.* (2006) Evidence for complete  
 487 denitrification in a benthic foraminifer. *Nature*, **443**, 93-96.

488 Rowe GT, Wei C, Nunnally C, *et al.* (in press) Comparative Structure and Dynamics of Food  
 489 Webs in the Deep Gulf of Mexico. *Deep-Sea Research II*.

490 Sen Gupta BK (1999) *Modern Foraminifera*. Kluwer Academic Publishers, Dordrecht, The  
 491 Netherlands.

492 Seibel BA, Walsh PJ (2003) Biological impacts of deep-sea carbon dioxide injection inferred  
 493 from indices of physiological performance. *Journal of Experimental Biology*, **206**, 641-  
 494 650.

495 Smith KL Jr, Baldwin RJ, Karl DM, Boetius A (2002) Benthic community responses to pulses in  
 496 pelagic food supply: North Pacific Subtropical Gyre. *Deep-Sea Research I*, **49**, 971-990.  
 497 Snider LJ, Burnett BR, Hessler RR (1984) The composition and distribution of meiofauna and  
 498 nanobiota in a central North Pacific deep-sea area. *Deep-Sea Research*, **31**, 1225-1249.  
 499 Szarek R, Nomaki H, Kitazato H (2007) Living deep-sea benthic foraminifera from the warm  
 500 and oxygen-depleted environment of the Sulu Sea. *Deep-Sea Research II*, **54**, 145-176.  
 501 Tamburri MN, Peltzer ET, Friederich GE, Aya I, Yamane K, Brewer PG (2000) A field study of  
 502 the effects of CO<sub>2</sub> ocean disposal on mobile deep-sea animals. *Marine Chemistry*, **72**,  
 503 95-101.  
 504 Thistle D, Sedlacek L, Carman KR, Fleeger JW, Brewer PG, Barry JP (2007) Exposure to carbon  
 505 dioxide-rich seawater is stressful for some deep-sea species: an in situ, behavioral study.  
 506 *Marine Ecology Progress Series*, **340**, 9-16.  
 507 van Oevelen D, Soetaert K, Middelburg JJ, Herman PJM, Moodley L, Hamels I, Moens T, Heip  
 508 CHR (2006) Carbon flows through a benthic food web: Integrating biomass, isotope and  
 509 tracer data. *Journal of Marine Research* **64**, 453-482.  
 510 Watanabe Y, Yamaguchi A, Ishidai H, *et al.* (2006) Lethality of increasing CO<sub>2</sub> levels on deep-  
 511 sea copepods in the western North Pacific. *Journal of Oceanography*, **62**, 185-196.

512 Table 1. Mean abundances of live foraminifera [(#·10cm<sup>3</sup>); listed for all foraminifera and by foraminiferal group] for each treatment,  
 513 presented by experiment. The number of replicates (# cores) is also listed. SD = Standard deviation.

514

515	Treatment	Experiment	# cores	Mean (SD)			
516				<u>Total</u>	<u>Thecate</u>	<u>Agglutinated</u>	<u>Calcareous</u>
517	Initial Control	1	3	4.8 (0.8)	0.9 (0.5)	1.6 (0.8)	2.3 (0)
518	Control	1	3	9.5 (2.9)	3.3 (2.8)	3.0 (0.8)	3.2 (1.8)
519		2	2	9.4 (1.8)	2.9 (3.3)	2.0 (0.9)	4.6 (0.5)
520	Elevated Center	1	3	8.3 (4.5)	2.4 (2.2)	3.0 (2.6)	2.9 (1.6)
521	Elevated Edge	1	2	14.3 (16.6)	6.5 (7.0)	6.1 (7.5)	1.7 (2.0)
522		2	2	15.6 (6.9)	1.4 (0.6)	3.6 (1.5)	10.5 (5.0)
523	Direct Contact	1	3	4.7 (1.7)	3.0 (0.9)	1.7 (1.4)	0
524		2	2	6.5 (7.4)	4.0 (4.6)	2.5 (2.8)	0

**Figure Legends.**

**Figure 1.** Schematic representation of experimental configurations (a. Experiment 1; b. Experiment 2). The CO<sub>2</sub> cylinders are shown as gray circles, and CTDs equipped with pH sensors are shown as squares. Pushcores analyzed for this study are shown as smaller circles; pushcore treatment categorizations are: open = Control; / = Initial Control, black = Direct Contact, + = Elevated Center, \* = Elevated Edge; ~ = DMSO.

**Figure 2.** Photographs of injector pushcores placed in the seafloor. a. Representative pushcore from Elevated Center treatment immediately after the ROV manipulator released its grip on the thick injector tubing (upper left). Inset: Note the stream of whitish fluid (i.e., CellTracker Green and DMSO) entering the header space from the capillary tubing, which extends through the corer top, approximately to the arrow. b. Suite of three Control pushcores incubating in situ during Experiment 1.

**Figure 3.** Reflected light (a, c, e, g) and corresponding epifluorescence (b, d, f, h) micrographs of calcareous foraminifera collected at experimental termination. Those labeled with \* were from a Direct Contact treatment core (935-JB2); those without \* were from a Control treatment core (937-JB7). Note that none of the Direct Contact specimens fluoresce brightly. a, b. Top row: *Quinqueloculina venusta*, bottom row: *Pyrgo murrhina*; c, d. *Cassidulina* cf. *delicata*; e, f (left to right): *Hoeglundina elegans*, *Cibicidoides* sp.; g, h. *Uvigerina canariensis*. Scale bars: 100 µm.

**Figure 4.** Reflected light (a, c, e, g) and corresponding epifluorescence (b, d, f, h) micrographs of agglutinated and thecate foraminifera collected at experimental termination. Those labeled with \* were from Direct Contact treatment cores (935-JB2 or 935-JB3); those without \* were from Control treatment cores (937-JB6 or 937-JB7). Note that the Direct Contact specimen in c,d fluoresces brightly; other Direct Contract specimens were deemed dead. a, b (left to right). Unidentified allogromiid, saccamminid; c, d (left to right). *Pelosina* sp., unidentified allogromiid; e, f. *Hormosinella guttifera*; g, h. *Veleroninoides wiesneri*. Scale bars: 100  $\mu$ m.

**Figure 5.** Histograms of CellTracker Green labeled benthic foraminifera, presented as total density (specimens $\cdot$ 10cm<sup>3</sup>) and by group and treatment for each Experiment. Error bars reflect standard error.

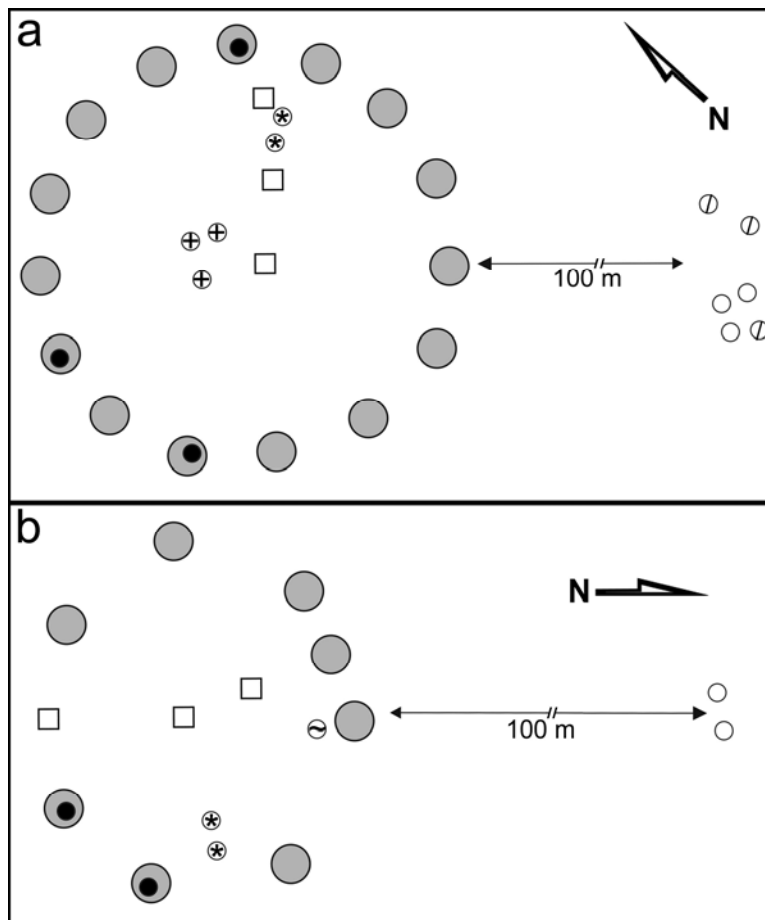


Figure 1.

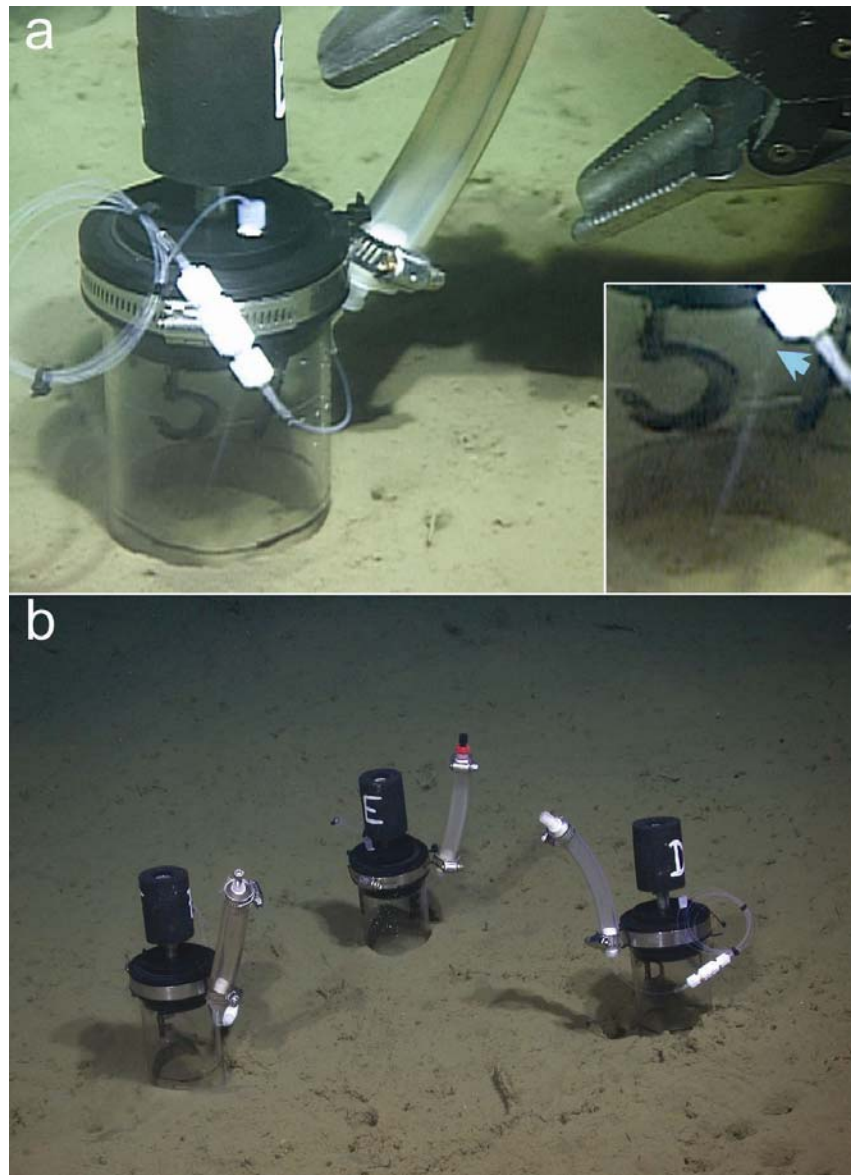


Figure 2.

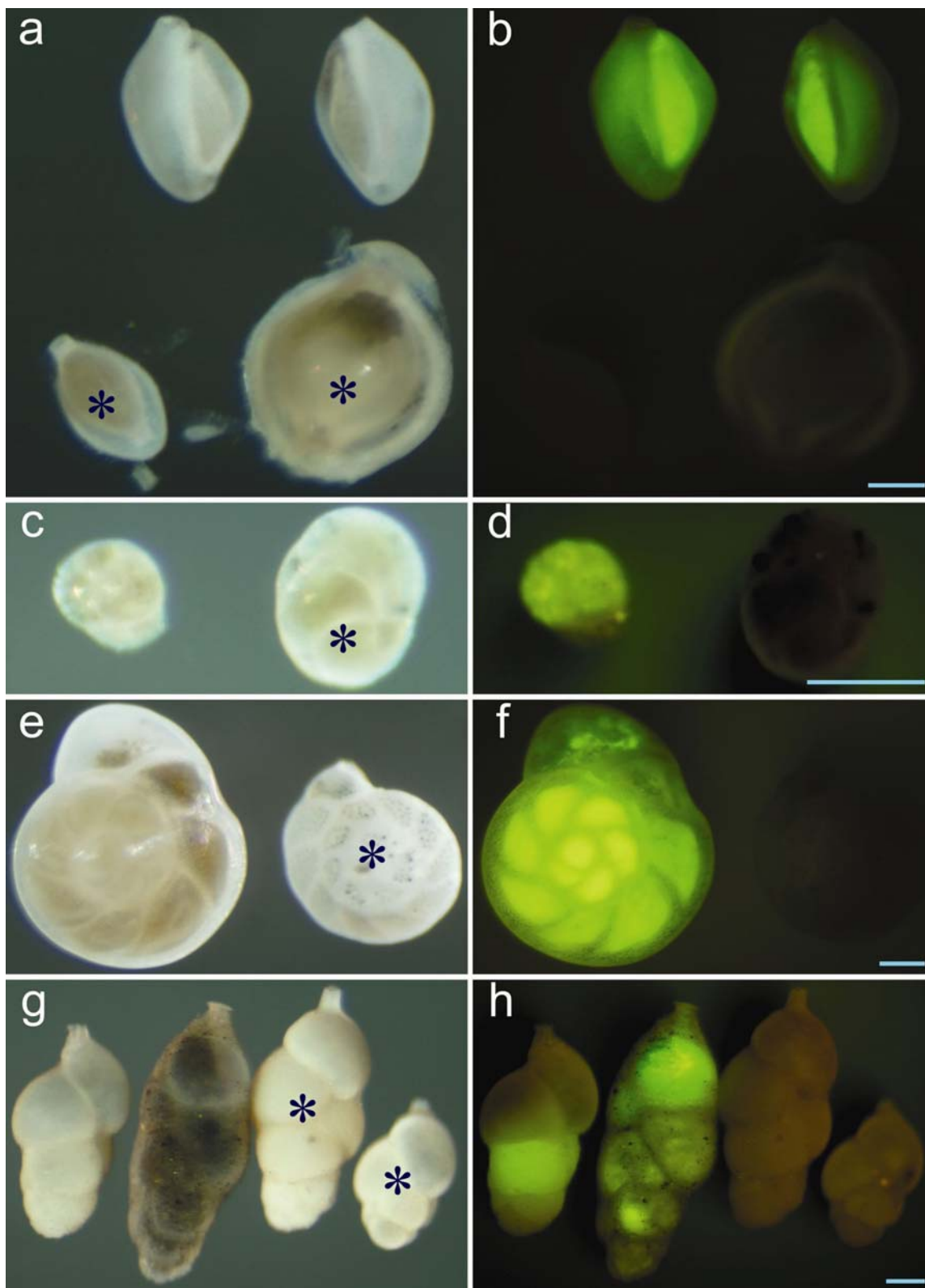


Figure 3.



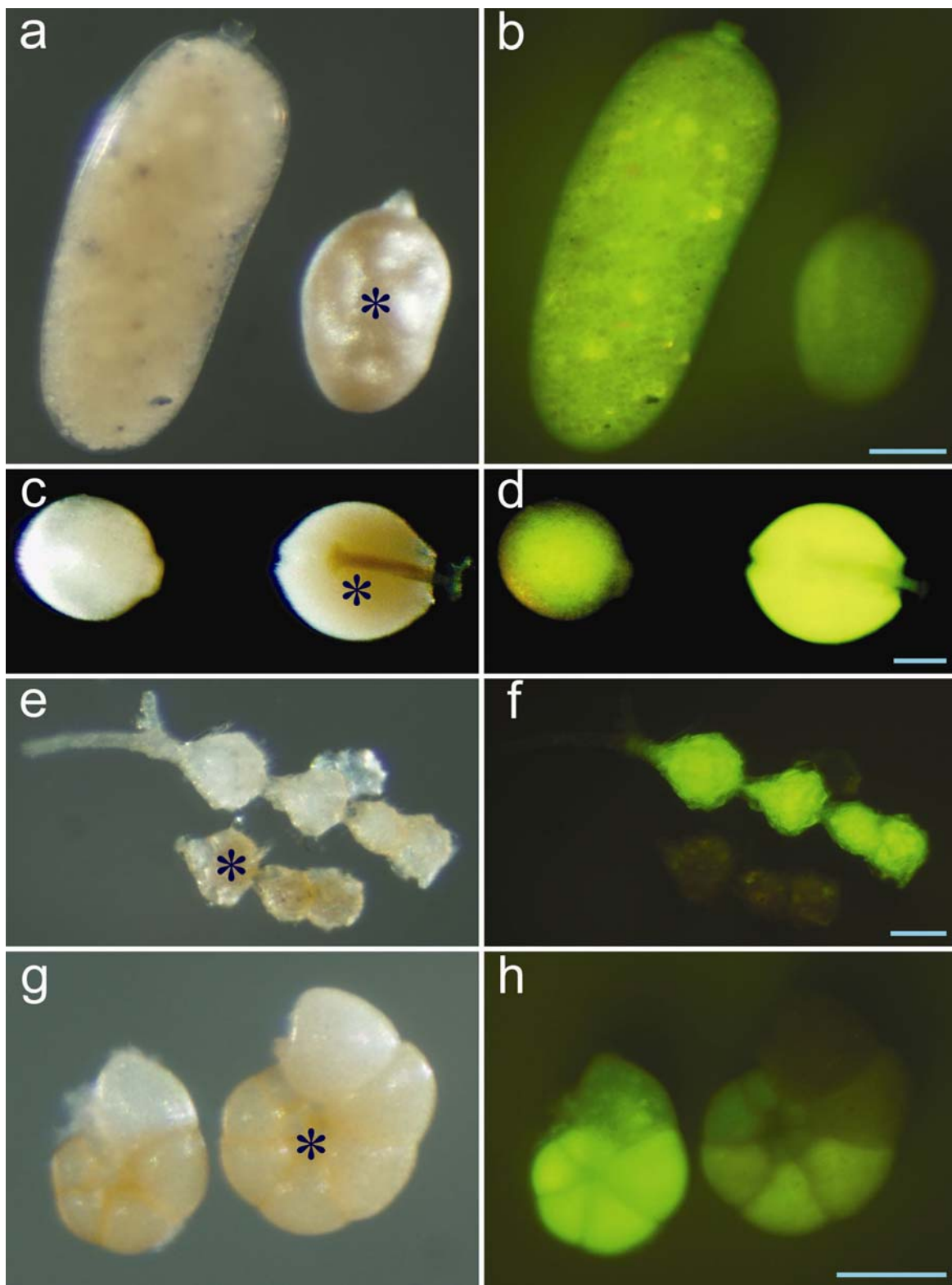


Figure 4.

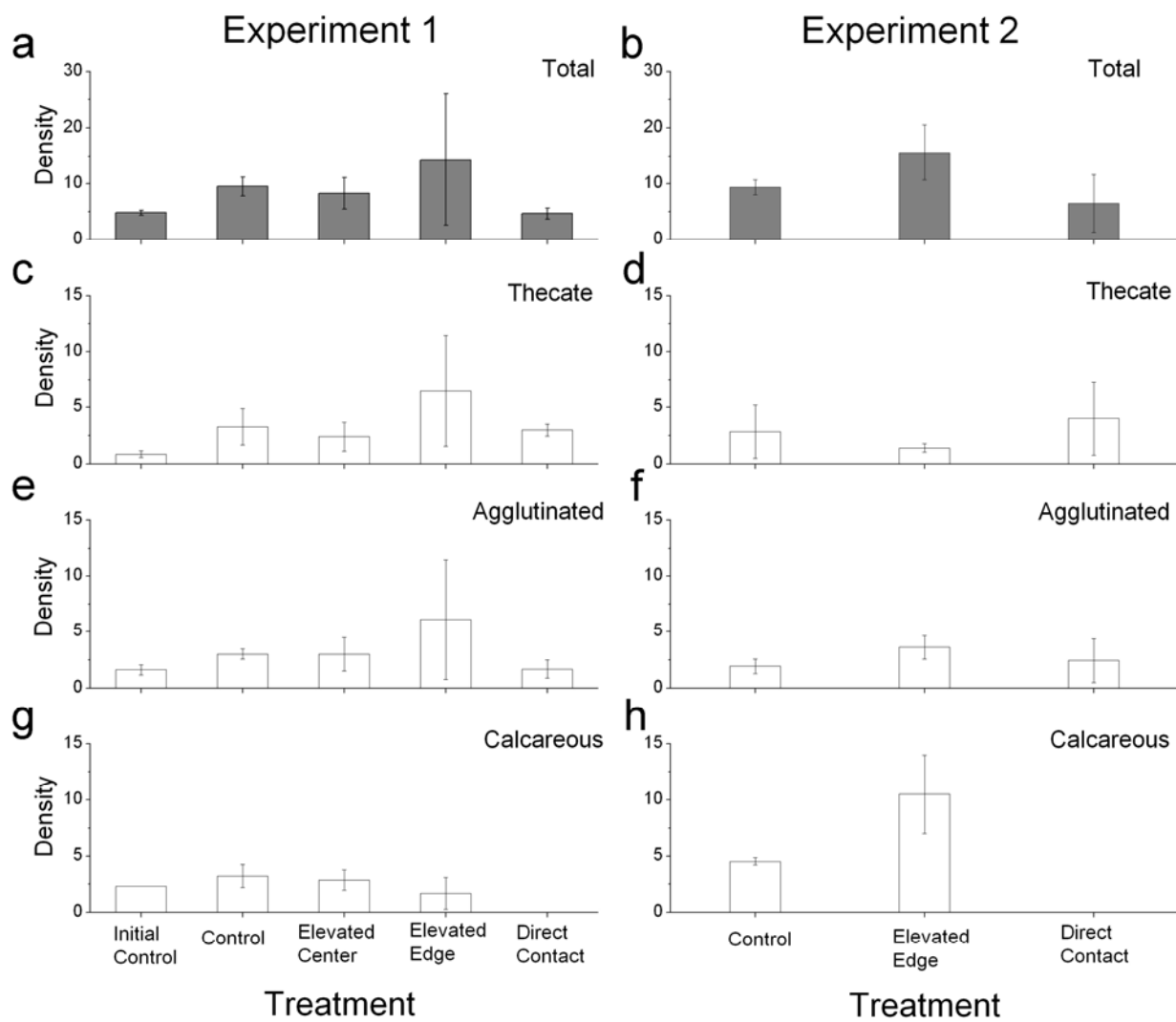


Figure 5.